

Source time function and time reversal technique

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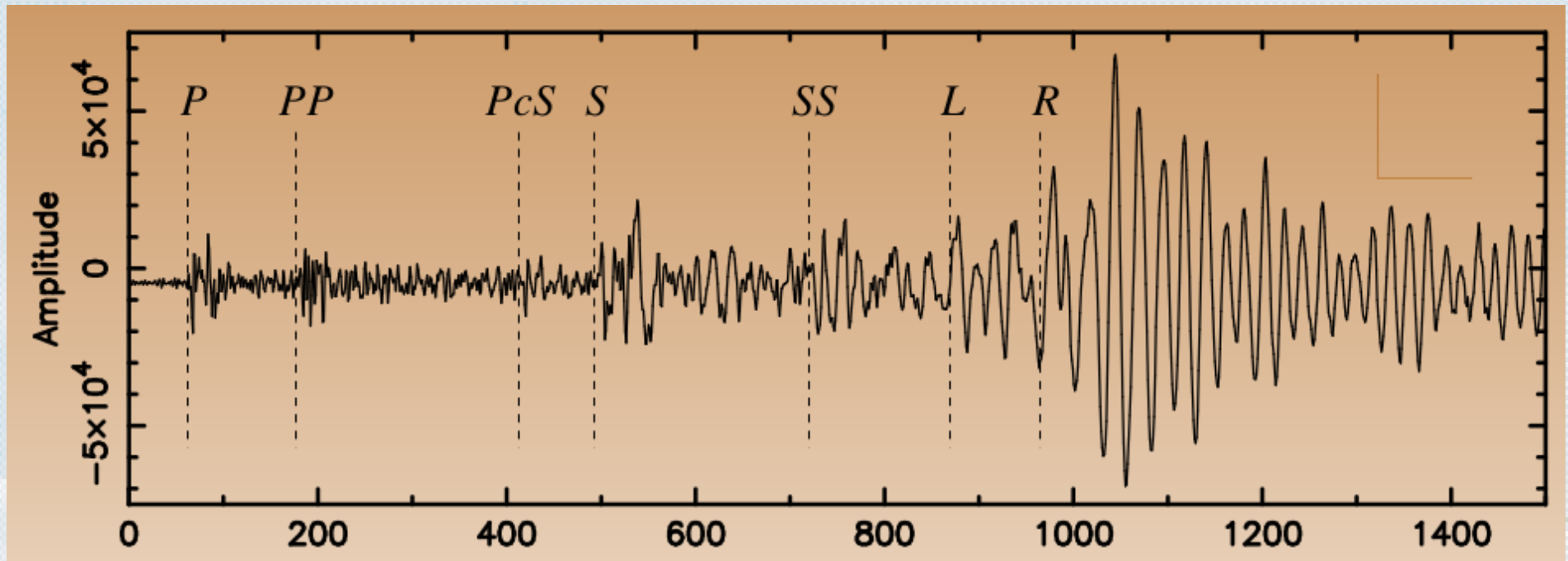
Seismological inference

Seismology deals with two broad classes of problems:

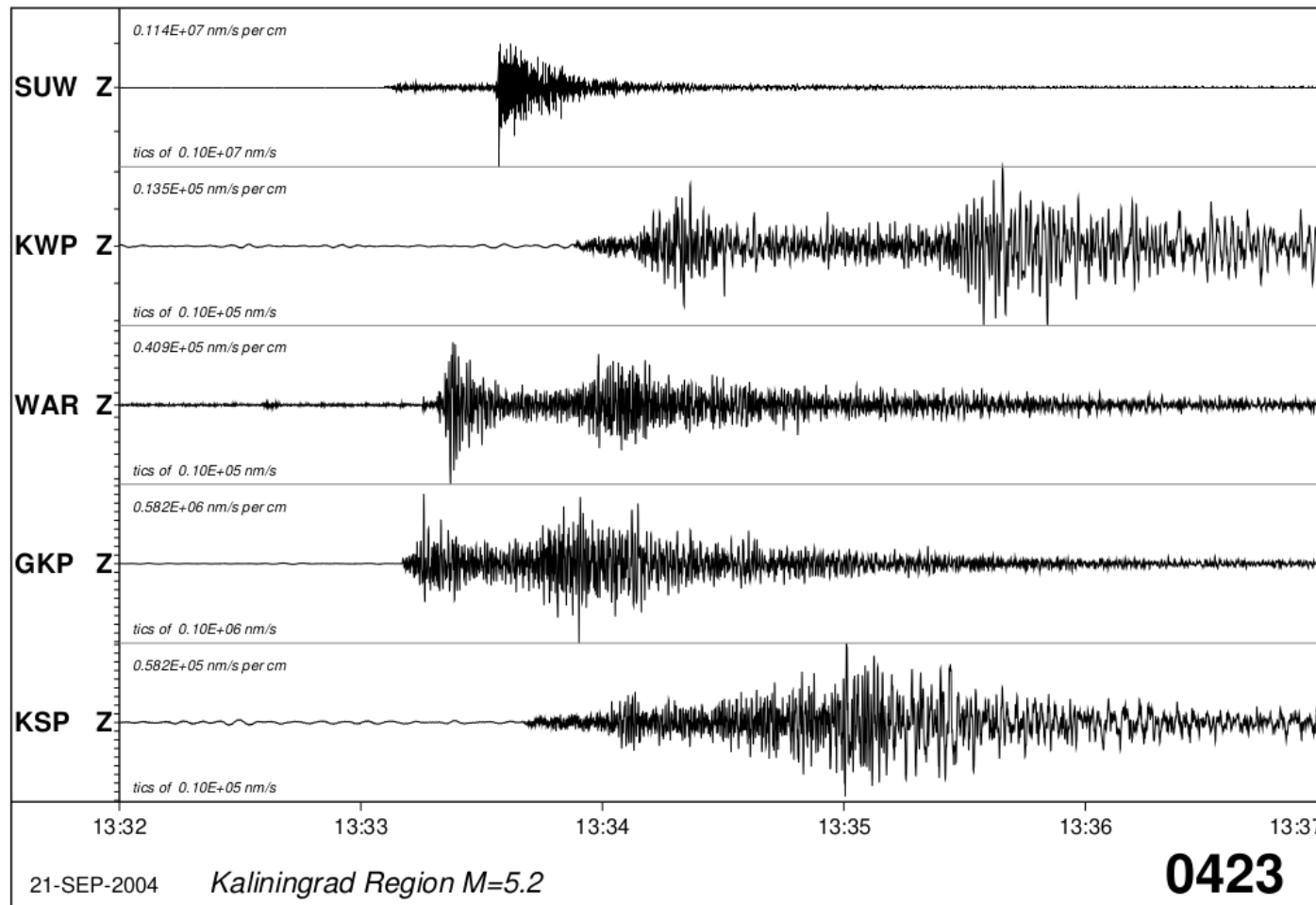
- ◆ Structure and composition of the Earth in various scales
global, regional (~ 2000 km), local (~ 100 km), etc.
- ◆ Earthquakes: their origins, physics, effects and (possible) prediction

Inference is always based on records made on surface (space)
There is no possibility of methodological direct measurements
(especially seismic sources).

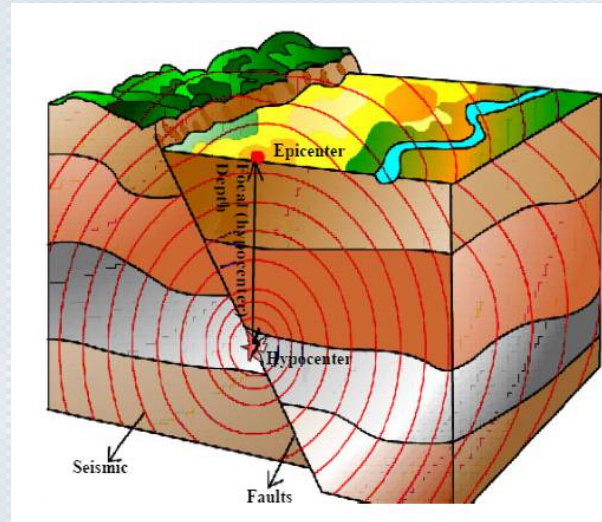
Seismic waves - primary source of information



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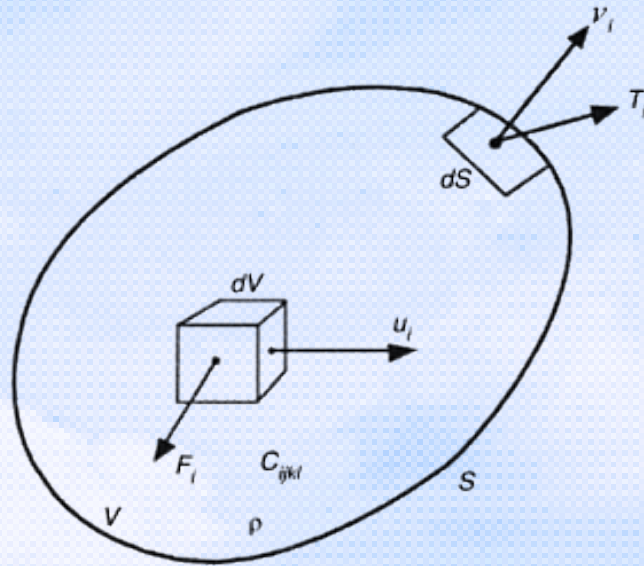


Seismic waves - information contents



$$\mathbf{u}(\mathbf{r}, t) = \int_{R', T'} G(\mathbf{r}, \mathbf{r}', t - t') \mathbf{S}(\mathbf{r}', t') d\mathbf{r}' dt'$$

Mathematics of elastic waves



$$T_i = \tau_{ij} n_j$$

$$\epsilon_{ij} = \frac{1}{2}(\partial_i u_j + \partial_j u_i)$$

$$\frac{d}{dt} \int_V \rho v_i(x, t) dV = \int_S T_i(x, t) dS + \int_V F_i(x, t) dV$$

Mathematics of elastic waves

Using Stock's theorem

$$\int_V \rho \frac{\partial^2 u_i}{\partial t^2} dV = \int_V \frac{\partial \tau_{ij}}{\partial x_j} dV + \int_V F_i(x, t) dV$$

$$\rho \ddot{u}_i = \partial_j \tau_{ij} + F_i$$

Homogeneous isotropic elastic medium

Hook's law:

$$\tau_{ij} = c^{ijkl} \epsilon_{kl}$$

Isotropic materials:

$$c^{ijkl} = \lambda \delta_{ij} \delta_{kl} + 2\mu \delta_{ik} \delta_{jl}$$

Homogeneous medium: $\lambda, \mu = \text{const.}$

$$\rho \ddot{\mathbf{u}} = (\lambda + \mu) \nabla (\nabla \cdot \mathbf{u}) + \mu \nabla^2 \mathbf{u} + \mathbf{F}$$

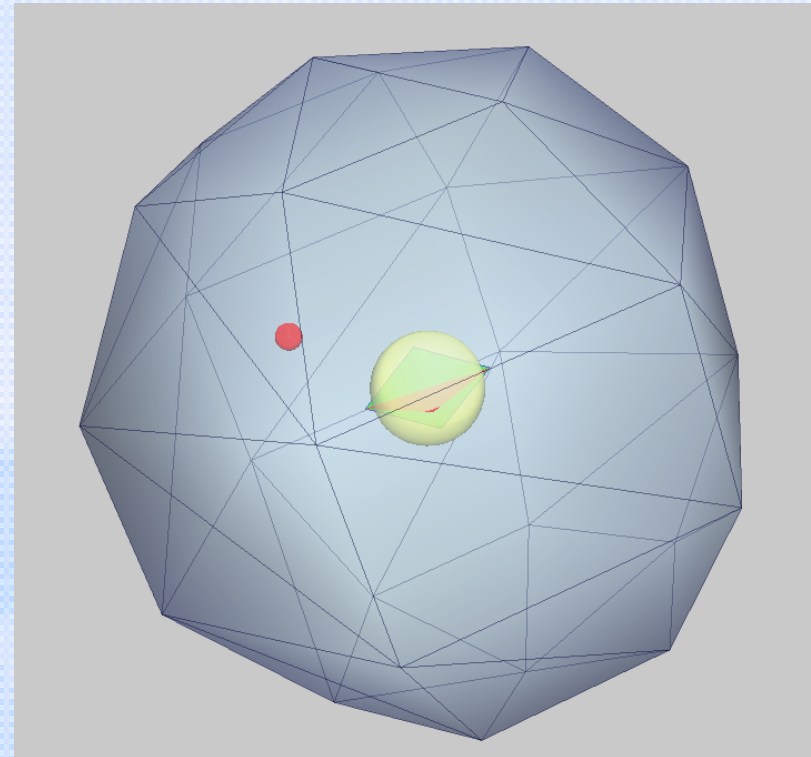
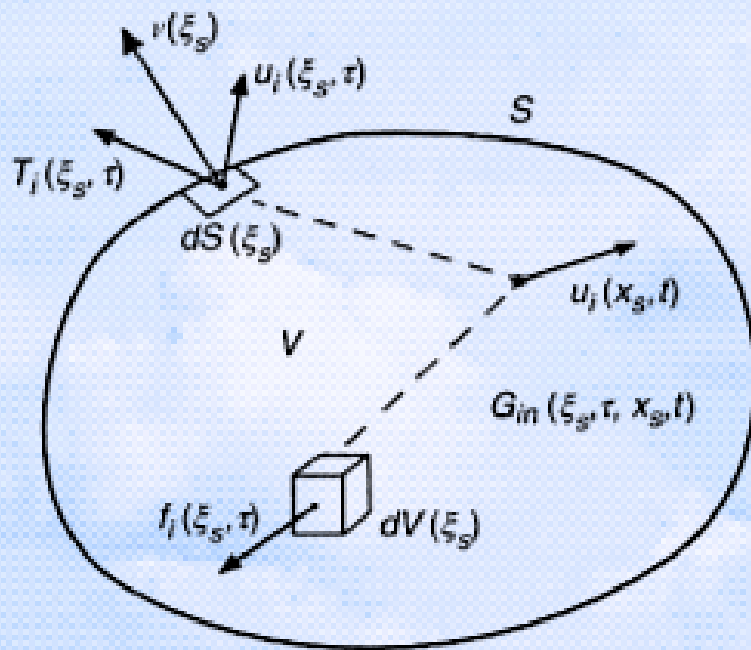
Representation theorem

Desired explicit solution for given source term (F_i)

Green's function: $(\rho\delta_{ik}\partial_t^2 + \partial_j\tau_{ij}) G_{ik} = e_k\delta(\mathbf{r}, \mathbf{t})$

$$\begin{aligned}
 u_n(x_s, t) &= \int_{-\infty}^{\infty} dt' \int_V \mathbf{f}_i(\xi, t') G_{ni}(\xi, t'; x_s, t) d\xi \\
 &+ \int_{-\infty}^{\infty} dt' \int_S G_{ni}(\xi, t'; x_s, t) \mathbf{T}_i(\xi, t') d\xi \\
 &- \int_{-\infty}^{\infty} dt' \int_S C_{ijkl} \mathbf{u}_i(\xi, t') G_{nk,l}(\xi, t'; x_s, t) n_j(\xi) d\xi
 \end{aligned}$$

Interpretation - forward problem



Given $f_i, T_{|\Sigma}, u_{|\Sigma}$ one can **uniquely** determine $u(\mathbf{x})$

Interpretation - inverse problem

Given $u(\mathbf{x}, t)$ can source be determined uniquely ?

NO!

Three **equivalent** class of source models:

1. f_i - body force (intuitive)
2. $u|_{\Sigma}$ - kinematic (dislocation source)
3. $T|_{\Sigma}$ - dynamic (stress drop, friction, etc.)

Equivalence: Body force model

Point source approximation: $L \ll \lambda, L \ll R$

$$u_n(x_r, t) = \int_{-\infty}^{\infty} f_i(x_s, t') G_{ni}(x_s, t'; x_r, t) dt'$$

For DC source system (4 elementary sources)

$$u_n(x_r, t) = (e_k e_j + e_j e_k) \int_{-\infty}^{\infty} M(t') \frac{\partial G_{nk}(x_s - x_r, t - t')}{\partial x_j} dt'$$

Equivalence: Kinematic source model

$$u_n(x_s, t) = \int_{-\infty}^{\infty} dt' \int_S C_{ijkl} \Delta u_i(\xi, t') G_{nk,l}(\xi, t'; x_s, t) n_j(\xi) d\xi$$

Equivalence: Kinematic source model

Assumptions:

- ◆ point source, homogeneous medium, non-rotating rupture
- ◆ $\mathbf{e}^s$ - slip unit vector ($\Delta \mathbf{u} = \mathbf{e}^s \Delta u$)
- ◆ $\mathbf{e}^n$ - vector normal to rupture plane
- ◆ $\mathbf{e}^n \cdot \mathbf{e}^s = 0$ - pure shearing

$$u_n(x_s, t) = \int_{-\infty}^{\infty} S \Delta u \left(\lambda e_l^s e_l^n \delta_{ij} + \mu (e_i^s e_j^n + e_j^s e_i^n) \right) G_{ni,j}(x_s, t'; x_r, t) dt'$$

Equivalence:

Kinematic model:

$$u_n(x_s, t) = \int_{-\infty}^{\infty} (e_k^s e_j^n + e_j^s e_k^n) \underbrace{\mu S \Delta u(t')}_{M_o(t')} G_{nk,j}(x_s, t'; x_r, t) dt'$$

Body force model:

$$u_n(x_r, t) = (e_k e_j + e_j e_k) \int_{-\infty}^{\infty} M(t') G_{nk,j}(x_s, t'; x_r, t) dt'$$

Double couple $\equiv$ shear dislocation source

Far field solution - homogeneous medium

Moment tensor

$$M_{ij}(t) = (e_k^s e_j^n + e_j^s e_k^n) \mu S \Delta u(t) = (e_k^s e_j^n + e_j^s e_k^n) M_o(t)$$

Source time function

$$M_{ij}(t) = m_{ij} M_o S(t)$$

Solution

$$u_i(x_s, t) = \frac{M_o}{4\pi\rho v_p^3 r} \gamma_i \gamma_j \gamma_k m_{jk} \dot{S}(t - r/v_p)$$

Moment tensor inversion

$$u_n(x_r, t) = \int_{-\infty}^{\infty} M_{ij}(t') G_{ni,j}(x_s, t'; x_r, t) dt'$$

$$M_{ij}(t) = (e_k^s e_j^n + e_j^s e_k^n) M_o S(t)$$

Moment tensor inversion should comprise

-
- ➔ Nm=1 Seismic scalar moment $M_o = \mu S \Delta u_o$
 - ➔ Nm=6 “*Fault plane solution*” $e_k^s e_j^n + e_j^s e_k^n$
 - ➔ Nm=(∞) Source time function $S(t) = \Delta \dot{u} / u_o$
-

Inversion seismograms for STF

◆ Standard approach (waveform inversion)

$$||u_{syn}(x_r, x_s, S(t)) - u_{obs}(x_r, x_s)|| + [\text{a priori}] = \min$$

◆ Empirical Green Function (Hartzell)

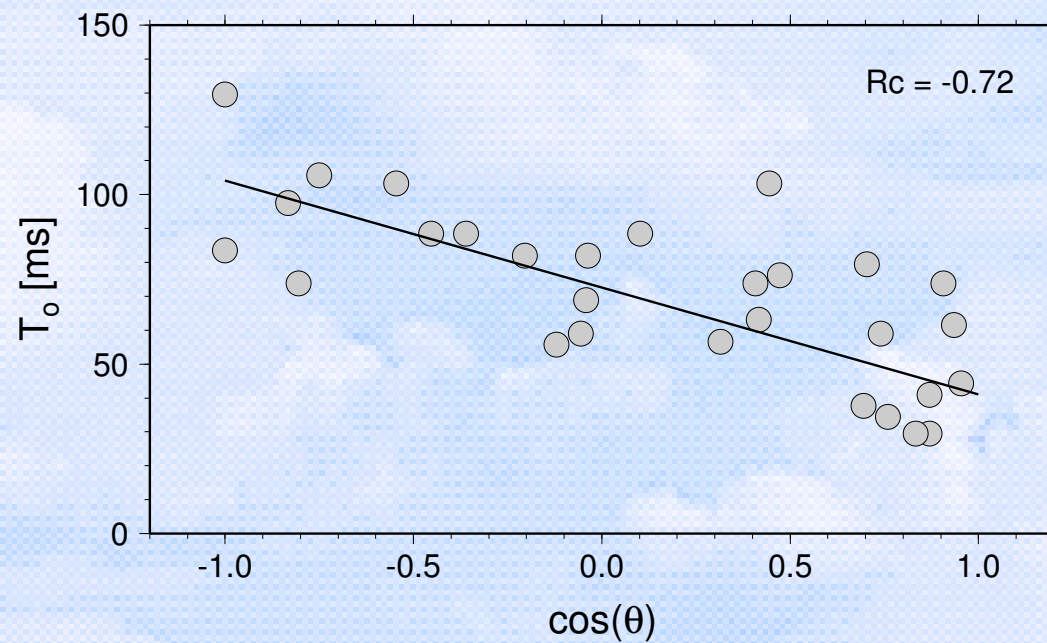
$$S(t) = \frac{u_{main}(t)}{u_{small}(t)} \quad (S_{small}(t) = \delta(t))$$

◆ Back projection (time reversal)

Physics inferred from STF

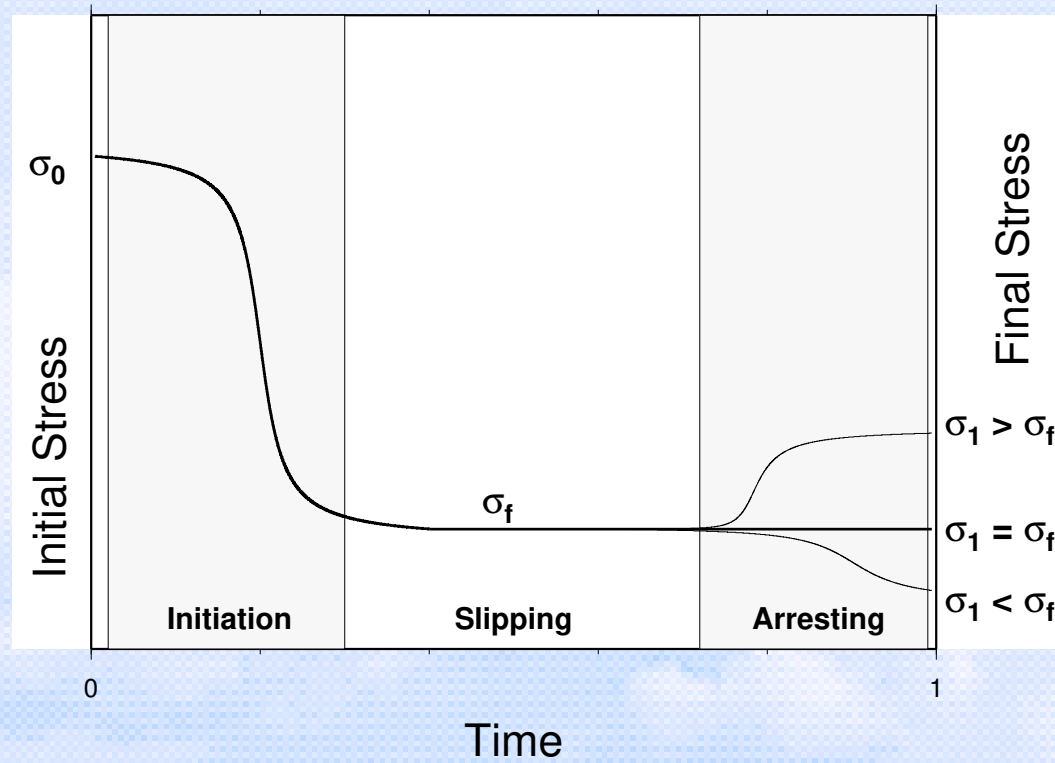
◆ rapture velocity (Ben-Menahem)

$$T_i = T_o - \delta T \cos(\theta)$$



Physics inferred from STF

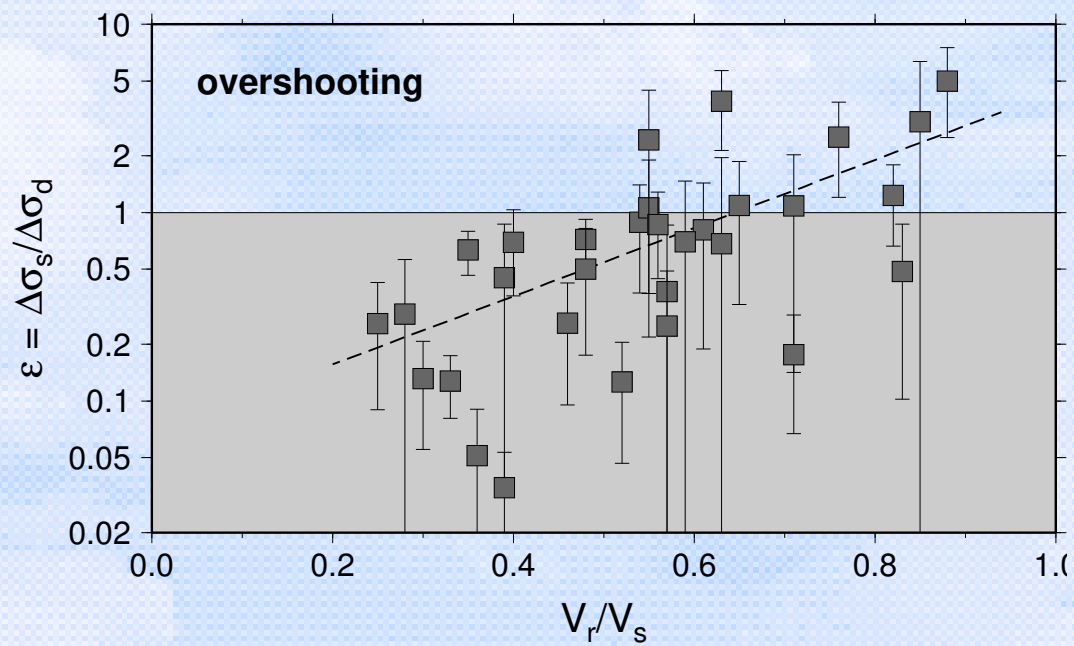
◆ arresting mechanisms



Physics inferred from STF

◆ dynamic stress drop

$$\Delta\sigma_d = \frac{M_o}{4\pi v_r^3} \xi \frac{\delta S}{\bar{S}}$$



Time reversal - Green's function

$$\frac{1}{c^2} \frac{\partial^2 G(x, t)}{\partial t^2} - \Delta G(x, t) = \delta(x - x_o, t - t_o)$$

Solution:

$$u(x, t) = \int_0^T dt' \int_{V_o} dx' G(t, x; t', x') S(x', t')$$

$G(x, t; , x', t')$ “propagates” information from source (x') to receiver (x)

Green's function cd.

Assuming time translation invariance
(energy conservation - Noether's theorem)

$$G(x, t; x', t') = G(x, x'; t - t')$$

Homogeneous medium
translational invariance

$$G(x, t; x', t') = G(x - x'; t - t')$$

Retarded and advanced Green's functions

Retarded (casual) GF:

$$G^+(x - x', t - t') = \frac{1}{4\pi ||x - x'||} \delta \left(t - t' - \frac{||x - x'||}{c} \right) \Theta(t - t')$$

Advanced (anti-casual) GF:

$$G^-(x - x', t - t') = \frac{1}{4\pi ||x - x'||} \delta \left(t - t' + \frac{||x - x'||}{c} \right) \Theta(t' - t)$$

$$G^+(x - x', t - t') = G^-(x - x', t' - t)$$

Time Reversal - step I

Radiation from point source (O)

$$S(x, t) = S(t)\delta(x - x_o)$$

Solution at given receiver (R)

$$u_R(x, t) = \int_{OR} G^+(x - x_o, t - t',) S(t') dt'$$

Time Reversal - step II

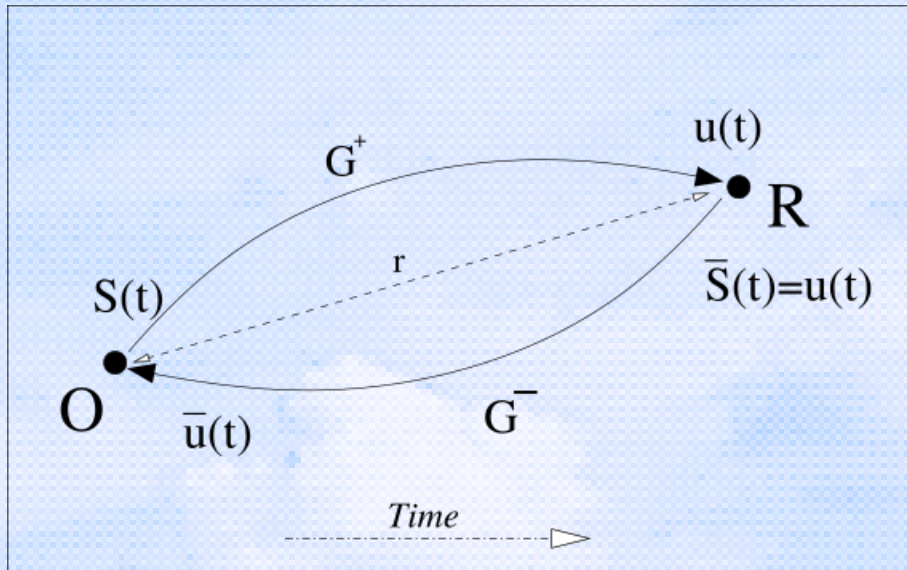
Anti-casual propagation from R to O

$$\bar{u}(x, t) = \int_{RO} G^-(x - x_R, t - t',) \bar{S}(t') dt'$$

From physical point of view it corresponds to recording at the point O incoming waves, for example, “refracted” at the point R

$$\bar{S}(t) = u_R(t)$$

Time Reversal Method (TRM) - basic principle



$$1. \quad u(r, t) = \int G^+(r, t - t') S(t') dt'$$

$$2. \quad \bar{S} = u(r, t)$$

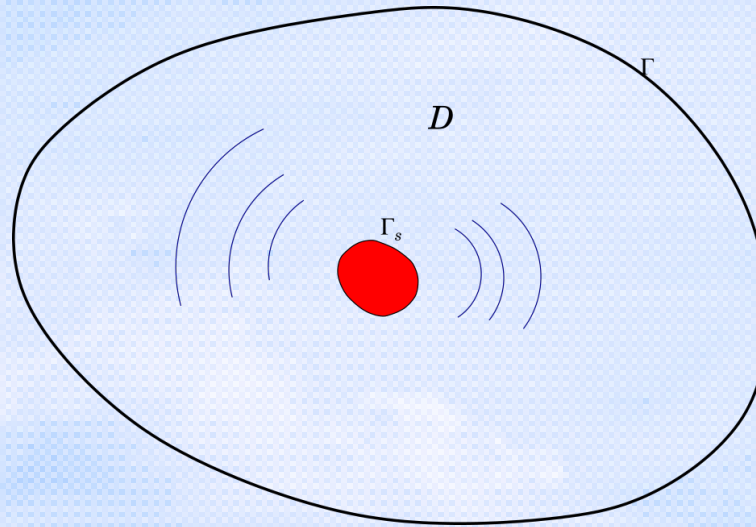
$$3. \quad \bar{u}(t) = \int G^-(r, t - t') \bar{S}(t') dt'$$

$$\bar{\mathbf{u}}(\mathbf{r}, \mathbf{t}) = \mathbf{k}(\mathbf{r}) \mathbf{S}(\mathbf{t})$$

However: $G^+(\cdot, t - t') = G^-(\cdot, -(t - t'))$

$$S(t) = \int G^+(t - t', r) u(-t') dt'$$

Time Reversal limitation - finite size source



$$u_D(x, t) = \int_0^T dt' \int_{V_o} dx' G^+(t, x; t', x') S(x', t')$$

$$u_D(x, t) = \int_{-\infty}^{\infty} dt' \int_{\partial\Gamma} ds' \left(u(x', -t') \frac{\partial}{\partial n'} G^+(x, t; x', t') - G^+(x, t; x', t') \frac{\partial}{\partial n'} u(x', -t') \right).$$

TRM as the inverse problem

$$\mathbf{m} = S(t)$$

$$\mathbf{d} = u(t)$$

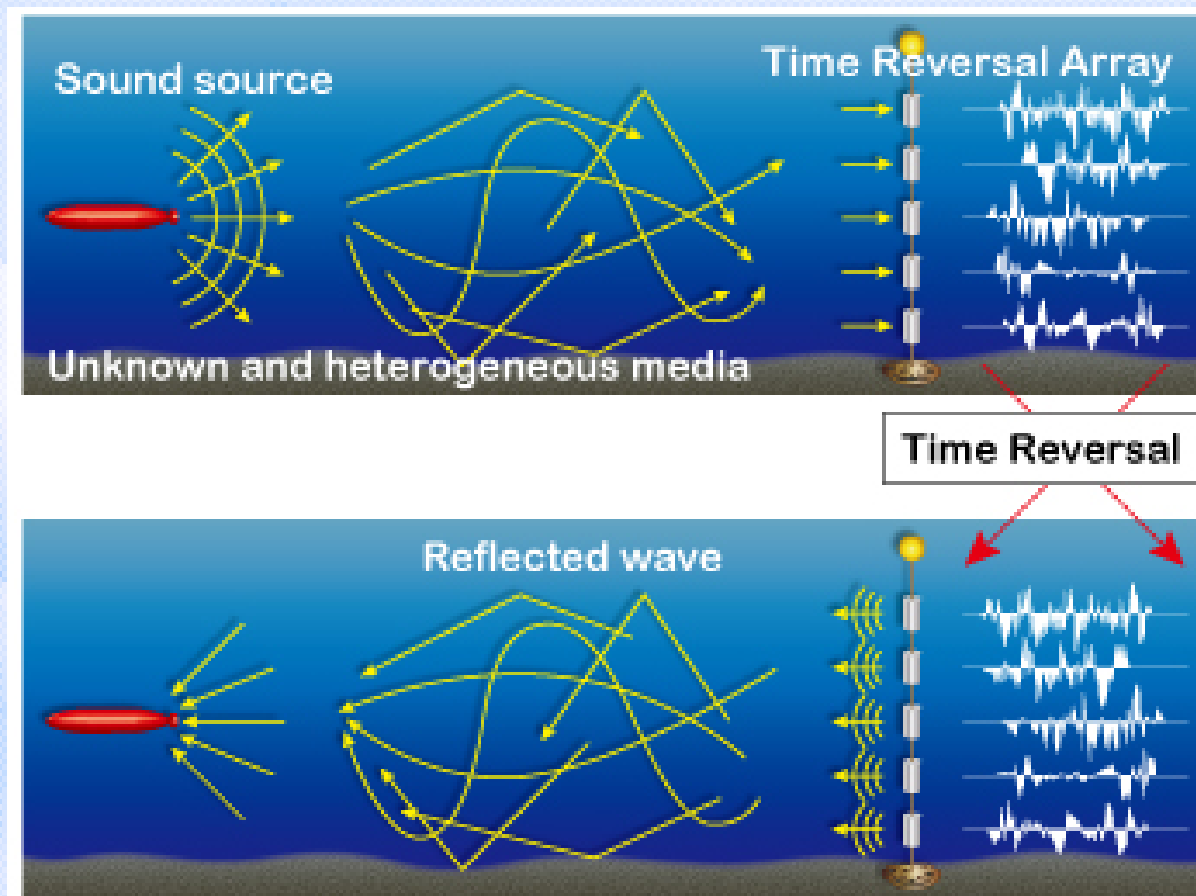
$$u(r, t) = \int G^+(r, t - t') S(t') dt' \quad \Rightarrow \quad \mathbf{d} = \mathcal{G} \mathbf{m}$$

$$S(t) = \int G^+(r, t - t') u(-t') dt' \quad \Rightarrow \quad \mathbf{m} = \mathcal{G}^{-1} \mathbf{d}$$

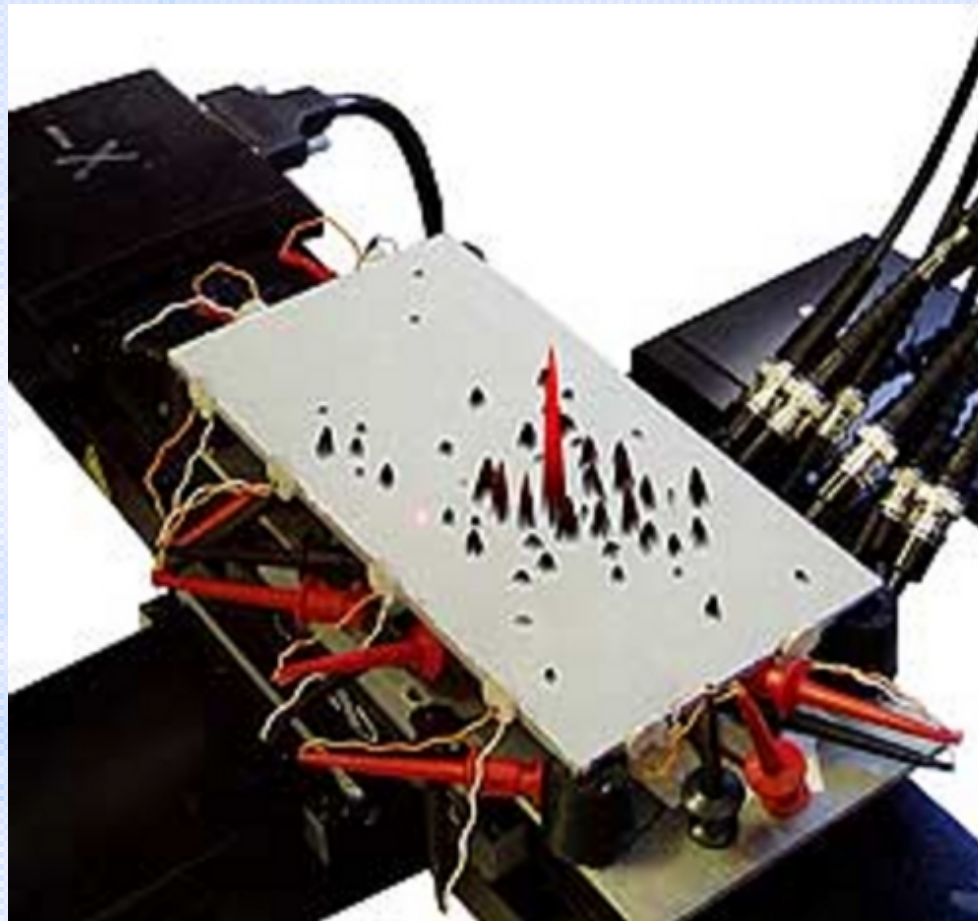
The method is computationally very efficient, but...

we need to know the medium: $G^+(\cdot, \cdot)$

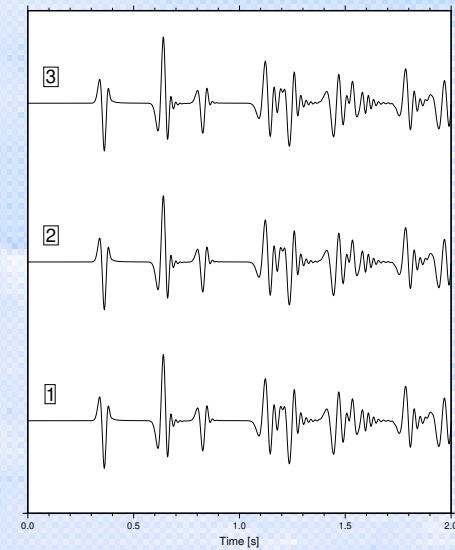
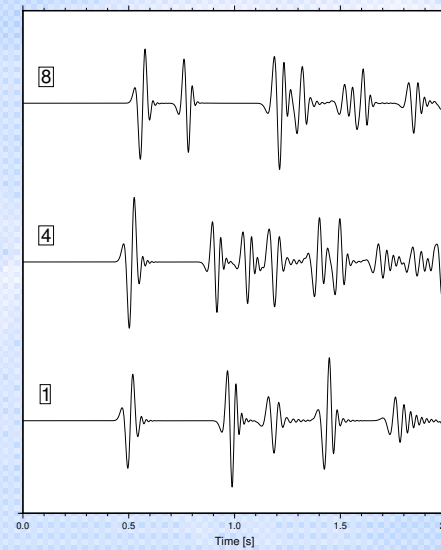
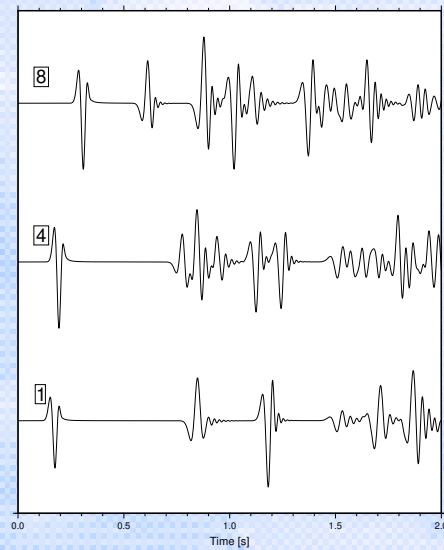
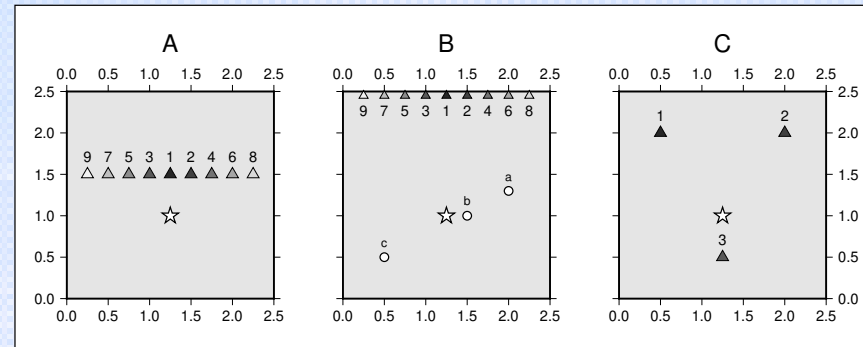
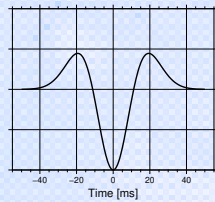
Physics of Time Reversal



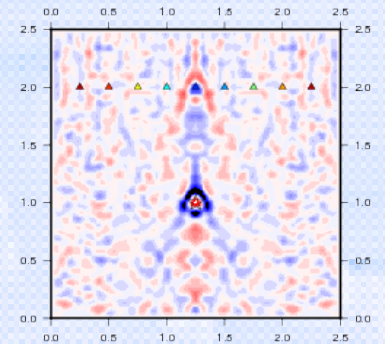
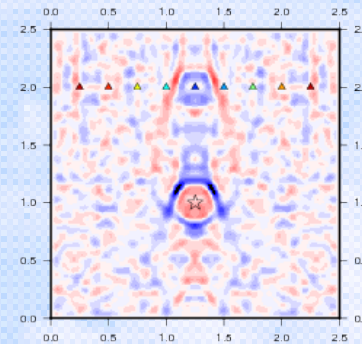
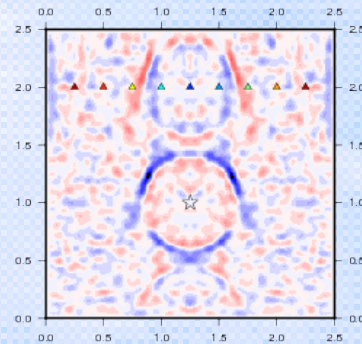
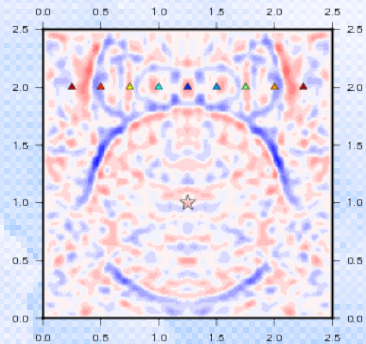
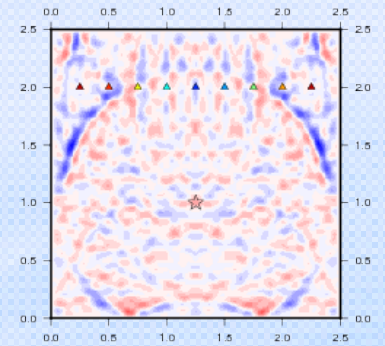
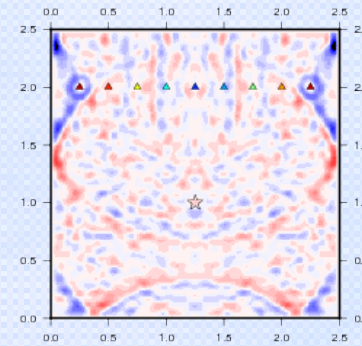
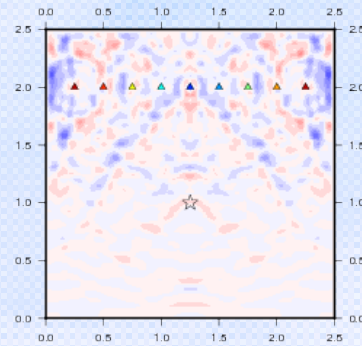
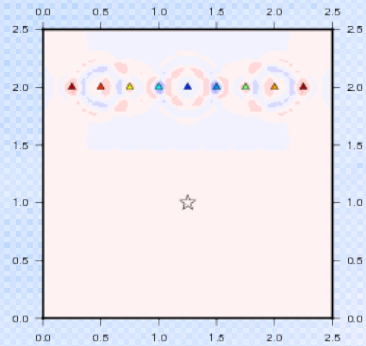
Experimental evidence (M. Fink)



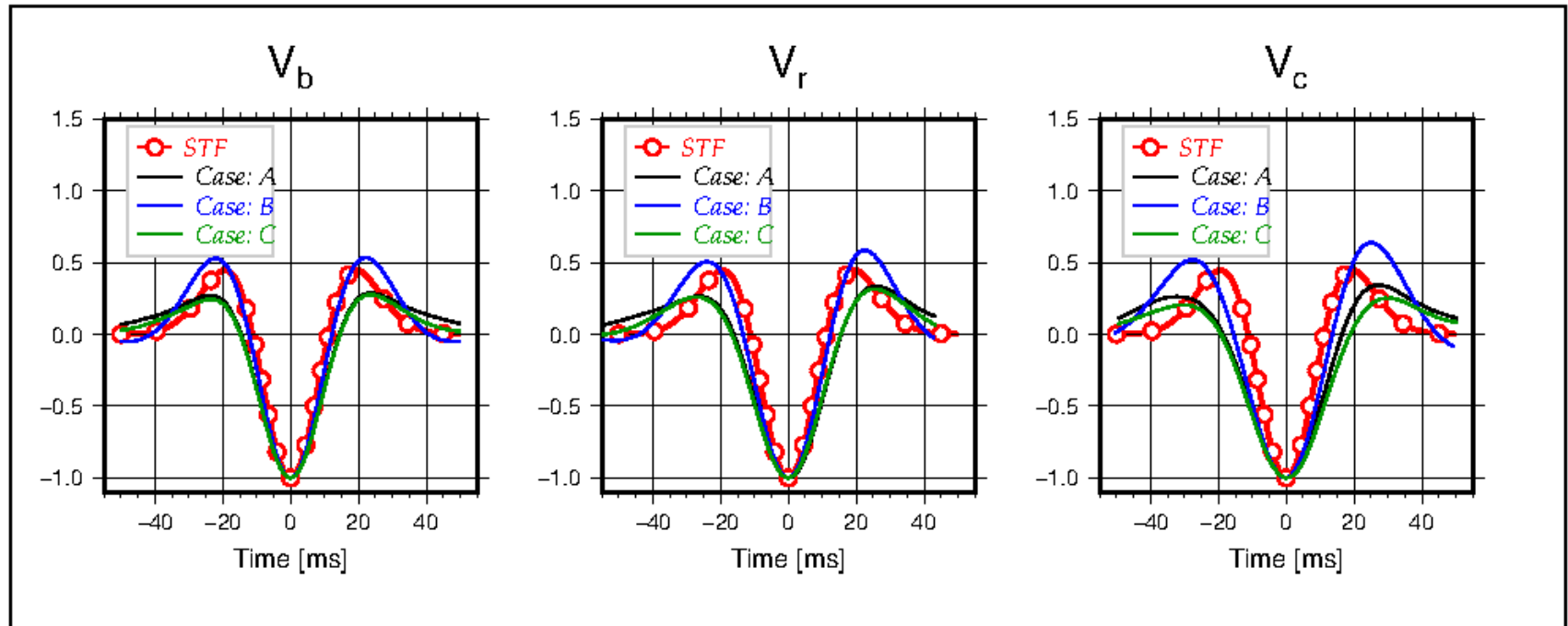
TRM and numerical simulations (K. Waskiewicz)



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Thank you